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2 **Sources of atmospheric volatile organic compounds during the Salt Lake regional**
3 **Smoke, Ozone and Aerosol Study (SAMOZA) 2022**

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13 **Key Points:**

- 14 • Methanol and ethanol dominated the total measured VOC abundance while isoprene
15 dominated total OH reactivity from VOCs in Salt Lake City
- 16 • Traffic and solvent use are roughly equivalent contributors to anthropogenic VOC
17 emissions in Salt Lake City during SAMOZA
- 18 • Ozone production was limited by VOCs in Salt Lake City in summer 2022

19 Abstract

20 We present measurements of volatile organic compounds (VOCs) and other trace gases taken in
21 Salt Lake City, Utah in August and September 2022. As part of the Salt Lake regional Smoke,
22 Ozone and Aerosol Study (SAMOZA), 35 VOCs were measured with two methods: a proton-
23 transfer-reaction time-of-flight mass spectrometer (PTR-ToF-MS) and 2,4-
24 dinitrophenylhydrazine (DNPH) cartridges analyzed by high-performance liquid
25 chromatography (HPLC). Over two months, the total measured VOCs averaged 32 ± 24 ppb
26 (mean \pm standard deviation) with the hourly maximum at 141 ppb, and the total calculated OH
27 reactivity averaged 3.7 ± 3.0 s⁻¹ (maximum at 20.7 s⁻¹). Among them, methanol and ethanol were
28 the most abundant VOCs, making up 42% of the ambient mixing ratio. Isoprene and
29 monoterpenes contributed 25% of the OH reactivity from VOCs, while formaldehyde and
30 acetaldehyde made up another 30%. The positive matrix factorization analysis showed 5 major
31 sources of VOCs, with 32% of abundance being attributed to secondary production/biogenic
32 sources, followed by 29% as traffic, 15% from industrial solvent use, and the rest from personal
33 care products (15%) and biomass burning (10%). Moderate smoke-impacted days on average
34 elevated various hazardous air pollutants (HAPs) by 45-217% compared to smoke-free days. The
35 ratio of OH reactivity from NO_x to that from VOCs showed that ozone production was mostly
36 VOC-limited throughout the campaign, consistent with our modeling study. VOCs and NO_x both
37 showed increased OH reactivity due to smoke influence. NO_x featured increased reactivity on
38 weekdays compared to weekends, an effect not shown for VOC reactivity during SAMOZA.

39 1 Introduction

40 Salt Lake City (SLC) metropolitan area in Wasatch Front, Utah, United States is a non-
41 attainment area for the 2015 National Ambient Air Quality Standards (NAAQS) for ozone (i.e.,
42 the 3-year average of the annual fourth-highest daily maximum 8-hour average ozone [MDA8]
43 not exceeding 70 ppb). With a total population of 2.13 million people, all 5 counties in the
44 Northern and South Wasatch Fronts were designated either marginal or moderate non-attainment
45 regions (https://www3.epa.gov/airquality/greenbook/anayo_ut.html). Ground-level ozone is a
46 photochemical product of nitrogen oxides (NO_x = NO + NO₂) and volatile organic compounds
47 (VOCs) in the presence of sunlight and has serious impacts on the health of the public,
48 ecosystem, and crops (Felzer et al., 2011; Nuvolone et al., 2018). However, its sources are
49 complex, including the local mix of anthropogenic origins of NO_x and VOCs, uncontrolled
50 sources such as biogenic or biomass burning, and long-range transport. Thus, ozone control
51 policy could vary city by city and require detailed knowledge of sources of precursors,
52 particularly for VOCs due to their complex origin in the urban atmosphere. Here we present
53 observations of a suite of VOCs and other trace gases made during the Salt Lake regional
54 Smoke, Ozone and Aerosol Study (SAMOZA) in August – September of 2022 and interpret the
55 data in terms of their origins and implications for ozone photochemistry in this region.

56 VOCs are emitted from a wide variety of sources in the urban atmosphere. Within cities,
57 common anthropogenic VOCs, such as aromatic compounds, often come from cars/trucks and
58 industry, with some contribution from long-range transport depending on the season and lifetime
59 of the compound (Hu et al. 2015a). Although vehicle design has been improved to reduce VOC
60 emissions, on-road emissions are still an important contributor to US urban VOCs (Warneke et
61 al., 2012). In the Intermountain West, emissions from oil and natural gas development have been
62 reported to enhance anthropogenic VOCs, including benzene and toluene, thus contributing to

63 high ozone abundance, for example, in the Colorado Front Range during spring and summer
64 (Abeleira et al., 2017). In Utah, the oil and gas production activities are mostly concentrated in
65 the east side of the State, i.e., the Uintah basin, and thus are thought to have no or little impact on
66 SLC (Womack et al., 2019). However, there are various petroleum industries including oil
67 refineries located in the Salt Lake City metropolitan area (Bhardwaj et al., 2021). Industrial sites
68 near urban areas can emit VOCs that are carried downwind to cities and contribute to pollution
69 (Gilman et al., 2013). In addition, Salt Lake City houses a variety of other manufacturing and
70 power plants in and around the city, from tubing and other materials manufacturing to aluminum
71 processing to power plants and refineries, which all emit VOCs that affect the area
72 (<https://deq.utah.gov/air-quality/2017-statewide-emissions-inventories>). Previous studies
73 reported abundances of methanol up to 80 ppb and acetaldehyde up to 7 ppb during wintertime
74 (Baasandorj et al., 2018), and formaldehyde reaching over 30 ppb throughout the year (Bhardwaj
75 et al., 2021) in SLC. These high levels of VOCs were thought to, at least to a certain extent, be
76 linked to such point sources from petroleum or other industries.

77 The use of solvents, also called volatile chemical products (VCPs), has become an
78 increasingly important source of VOCs, particularly as vehicle emissions have decreased in the
79 past few decades (McDonald et al., 2018). These VCPs also come from many of the products
80 that are used on a daily basis, such as D5 siloxane from personal care products and monoterpenes
81 from fragrances. Buildings themselves can emit VOCs from the offgassing of paint and wood
82 products, though the fluxes of such emissions are highly uncertain (Gkatzelis et al., 2021). Such
83 paints and coatings, personal care products, and cleaning products are thought to contribute the
84 most to VCP emissions, which could account for more than 50% of anthropogenic VOC
85 contribution to ozone formation in certain US cities (Coggon et al., 2020; Seltzer et al., 2021).

86 The relative contribution of VCPs to anthropogenic VOC emissions is likely to vary in
87 different urban areas, in part due to the large differences in population density. For example,
88 limited observational evidence suggested VCP emissions contributed to ~40% of anthropogenic
89 VOC emissions in Boulder, CO, up to half in Los Angeles, CA, but ~80% in New York City,
90 NY, while the remainder was often attributed to mobile emissions (60%, 50%, and 20%) in the
91 above 3 urban areas (Gkatzelis et al., 2021; McDonald et al., 2018). It is broadly consistent with
92 the recent bottom-up emission inventory suggesting that VCPs have become a dominant source
93 of anthropogenic VOCs in many US urban areas (Seltzer et al., 2021).

94 In Salt Lake City, the 2017 Northern Wasatch Front VOC and NO_x Inventories, which is
95 based on the EPA's 2017 National Emission Inventory (NEI), estimate that 44% of summertime
96 VOC emissions come from VCP/solvent uses, while only 23% of emissions come from onroad
97 sources (https://home.chpc.utah.edu/~u0864163/OZONE_public/NWF-SMOKE-Summary-Report.html). On an annual basis, the updated 2020 National Emissions Inventory from the EPA
99 also shows the importance of solvent use, with an estimated 27% of anthropogenic emissions
100 coming from solvent use, compared to 25% from onroad sources. The remaining anthropogenic
101 emissions come from other fuel combustion, such as aircraft and residential sources, as well as
102 industrial processes and agriculture.

103 Emissions of biogenic VOCs (BVOCs), including isoprene and monoterpenes (Folberth
104 et al., 2006; Guenther et al., 2012), are known to be important contributors to ozone formation
105 when interacting with urban NO_x sources because of their reactivity and emission fluxes, which
106 are often higher than those of typical anthropogenic VOCs (NRC, 1991). Isoprene concentration
107 is thought to be low in the western US compared to the east coast and southeast US, while

108 emissions of monoterpenes from needleleaf trees are supposed to be the dominant BVOCs in the
109 western US. However, uncertainties of these biogenic emissions are high, likely on the order of a
110 factor of 2 or higher (Sakulyanontvittaya et al., 2008). In addition, recent studies also suggested
111 that isoprene and monoterpenes in the urban atmosphere may in part reflect anthropogenic
112 emissions from VCPs or instrument interferences (Coggon et al., 2021; Gkatzelis et al., 2021; Hu
113 et al., 2015b; Peng et al., 2022). BVOC emissions are affected greatly by environmental
114 parameters including temperature and sunlight and are dependent on tree species or plant
115 functional types (Churkina et al., 2017; Gu et al., 2021; Ma et al., 2022). Salt Lake City boasts
116 around 85,000 trees within its limits (<https://www.slc.gov/parks/urban-forestry/>), which have
117 many benefits for the residents, but can also contribute to BVOC levels in the city. BVOCs can
118 be transported from upwind forests, which have been found to affect nocturnal chemistry and
119 contribute to high-ozone events the following day (Millet et al., 2016). These BVOC emissions
120 can complicate source attribution, as many compounds are co-emitted from anthropogenic and
121 biogenic sources including OVOCs such as methanol and acetone.

122 Smoke from wildland fires is another natural source that increases VOC levels and ozone
123 production in urban areas, especially in the western U.S. (Jaffe et al., 2020). Wildfires have been
124 suggested to be the second largest VOC source in the western US, though the uncertainty of their
125 emissions is also high and is thought to be on the order of a factor of 2-3 at least for this region
126 (Jin et al., 2023). The species emitted by wildfires depend on the biomass types and burn
127 conditions, among other factors (Gilman et al., 2015; Sekimoto et al., 2018) and can include
128 gases like formaldehyde and methanol, as well as NO_x (Liu et al., 2017; Permar et al., 2021). As
129 gases are transported from the fire, they undergo oxidation reactions via sunlight or chemical
130 species like OH radicals. The distance a smoke plume travels will affect the level of oxidation it
131 undergoes and will therefore dictate the composition of the plume as it reaches an urban site,
132 however in general smoke from wildland fires tends to be VOC-rich and NO_x -poor (Liang et al.,
133 2022; Permar et al., 2023). Thus each smoke event could be unique because of this dependence
134 on the location of emitted VOCs and those that reach the urban site. Once wildfire smoke
135 reaches a city, it mixes with urban emissions, often NO_x -rich, to degrade air quality and increase
136 ozone production. In many cases, ozone increases due to increased VOC levels (Dreessen et al.,
137 2016; Jaffe et al., 2013; Lill et al., 2022; Permar et al., 2023). Salt Lake City was seen to have
138 elevated ozone concentrations due to wildfires upwind of the region in the summer of 2015
139 (Gong et al., 2017; Horel et al., 2016). The chemical regime within the urban atmosphere may
140 shift as gas concentrations change, often transitioning from VOC-limited to transitional or NO_x -
141 limited scenarios as VOC levels rise without a corresponding increase in NO_x levels (Liang et
142 al., 2022; Rickly et al., 2023).

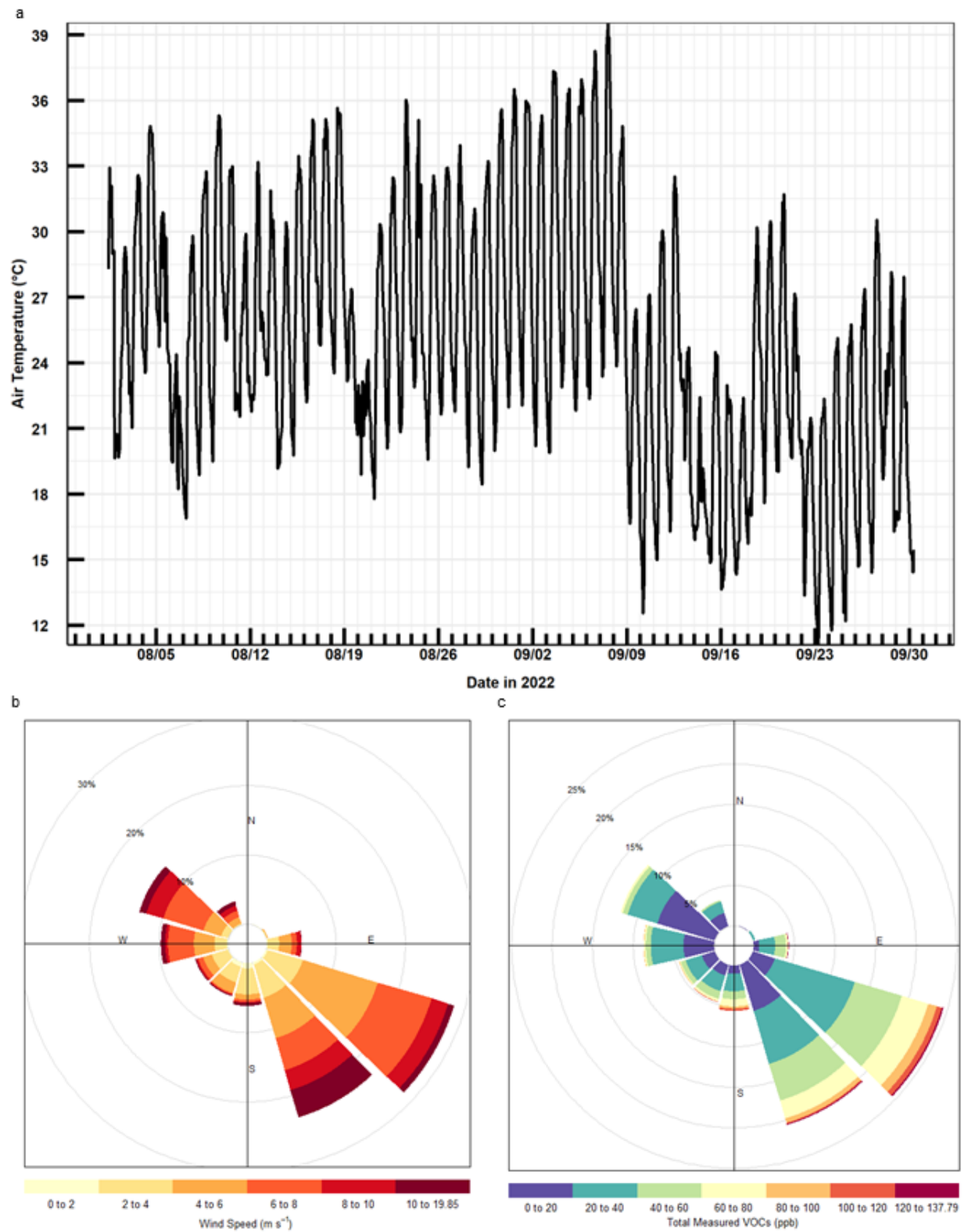
143 The SAMOZA field campaign aimed to better understand the factors leading to the high
144 levels of ozone in Salt Lake City by exploring its sources and how atmospheric composition is
145 affected by wildfire smoke (Jaffe et al., 2024). In this work, we report the abundance of 35
146 VOCs measured during the campaign and investigate their sources and atmospheric impact in
147 SLC. The abundance, composition, and reactivity of total and speciated VOCs in SLC are
148 presented. Their sources are then characterized using positive matrix factorization (PMF) and
149 corroborated by source tracer analysis. The impact of wildfire smoke is assessed for VOC
150 abundance and reactivity. Finally, the implications for photochemistry leading to ozone
151 production in SLC are explored. Further analysis and results from the campaign have also been
152 recently published. Jaffe et al. (2024) provides a general overview of the key results of the
153 campaign. Ninneman et al. (2023) presents results from a photochemical box model constrained

154 by observations from the campaign to describe ozone formation in the region. Lee and Jaffe
155 (2024) uses a generalized additive model to investigate the impact of wildfire smoke on ozone
156 concentrations in SLC over 16 years, including the measurement period of SAMOZA.

157 **2 Materials and Methods**

158 2.1 Measurement Site and Meteorological Data

159 Measurements for the SAMOZA study were conducted in Salt Lake City, Utah from
160 August 1 to September 30, 2022. The campaign measurements took place at the Utah
161 Department of Environmental Quality Technical Support Center (UDAQ UTC), located
162 approximately ~3 km east of the Salt Lake City International Airport at 40.778° N, -111.946° W.
163 UDAQ UTC sits on the northeast corner within a multi-agency business complex, surrounded by
164 the residential neighborhood to its north and east, and is 5 km away from SLC downtown. Being
165 at the intersection of 3 major Interstate Highways (~0.3 km east of I-215, 1.2 km north of I-80,
166 and 3 km west of I-15,) UDAQ UTC is expected to be a high NO_x and receptor site of diverse
167 urban sources.



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Figure 1. Meteorological conditions at the UDAQ UTC site during the SAMOZA field

170 campaign (August 1-September 30, 2022). a) air temperature ($^{\circ}\text{C}$); b) Wind direction as a
171 function of wind speed (m/s); and c) Wind direction as a function of total measured VOCs (ppb).

172
173 Figure 1 shows the air temperature measured at UDAQ UTC during the SAMOZA
174 campaign, as well as wind rose plots showing the wind direction as a function of wind speed and
175 total measured VOCs. Temperature data was collected by the UDAQ using an electronic thin
176 film air temperature sensor. Wind direction and speed were measured with 2D-ultrasonic
177 anemometer transducers. All meteorological instruments were situated on a tower with a height
178 of 10 meters above the roof at UDAQ UTC (Utah DAQ, 2022). The air temperature was warmer
179 for the first part of the campaign through September 9th ($27 \pm 5^{\circ}\text{C}$; mean \pm standard deviation).
180 Most days during the first period reached a daily highest temperature of over 30°C , with the
181 daily lowest temperature often being over 20°C ($21 \pm 2^{\circ}\text{C}$). From September 9th until the end of
182 the campaign, daily average temperatures dropped to $21 \pm 5^{\circ}\text{C}$. The daily low temperature during
183 this period dropped by nearly 6°C compared to the previous period ($15 \pm 2^{\circ}\text{C}$), with the high
184 temperature only exceeding 30°C six days out of the 21-day period. Wind came mostly from the
185 southeast at night, while during daytime wind direction tended to shift to the west or northwest
186 during SAMOZA (Figure S1). The highest wind speed seen was 19 m/s, while the average was
187 around 6 m/s.

188 2.2 VOC Measurements by PTR-ToF-MS

189 Ambient VOCs were measured using proton-transfer-reaction time-of-flight mass
190 spectrometry (PTR-ToF-MS 4000, Ionicon Analytik GmbH, Innsbruck, Austria). The conditions
191 in the drift tube were held constant during the campaign at 3.00 mbar, 60°C , and 815 V, which
192 made for an electric field (E/N) of 135 Townsend (Td). The PTR was located on the second floor
193 of UDAQ UTC. The ~ 10 m long sampling inlet was made from perfluoroalkoxy (PFA, 1/4''
194 OD) tubing and was situated on the roof of the building, ~ 20 m above ground level. The air was
195 subsampled by the PTR-ToF-MS through ~ 100 cm of 1/16'' (1.59 mm) OD polyetheretherketone
196 (PEEK) tubing maintained at 60°C . Ions from m/z 19 to 400 were measured once every minute.
197 Instrument background was taken approximately every 2½ hours by measuring VOC-free air
198 generated by ambient air passing through a heated catalytic converter (375°C , platinum beads, 1
199 wt% Pt: Sigma Aldrich).

200 In this work, we include 27 VOC species measured by PTR-ToF-MS in the analysis, which
201 are listed in the supplement (Table S1). Among them, calibrations were performed for 22 species
202 in two compressed gas standard cylinders following previously established procedures (stated
203 accuracy 5% at ~ 1 ppmv; Apel-Riemer Environmental, Inc., Miami, FL, USA; Permar et al.,
204 2021; Selimovic et al., 2022). One standard gas cylinder with 10 species was used every other
205 day during the campaign. A second standard gas cylinder containing 15 species was used every
206 other day for the first three weeks of the campaign. For isomers calibrated at the same m/z (i.e.,
207 methyl vinyl ketone and methacrolein at m/z 71.049; ethylbenzene and o-xylene at m/z 107.086;
208 1,2,4- and 1,3,5-trimethylbenzene at m/z 121.101), a weighted average sensitivity was found
209 using corresponding isomeric contributions per previous urban studies. Six-point calibrations
210 were performed by diluting the gas standards with the VOC-free air described above, at a range
211 between 1 and 7 ppb. Only calibrations with an r^2 above 0.998 and those sensitivities for the
212 same species that did not vary by 10% of the mean during the campaign were used. In addition,
213 D5 Siloxane was calibrated with a gas standard in June 2022 before the campaign (stated
214 accuracy 5% at ~ 1 ppm; Apel-Riemer Environmental, Inc., Miami, FL) using dynamic dilution

215 as described above. From quadrature addition of individual errors including calibrations and
216 mass flow controllers in the instrument, uncertainty for these species is estimated to be <15%.

217 Formaldehyde was calibrated after the campaign using a gas standard cylinder (stated
218 accuracy 5% at ~2 ppm; Airgas USA LLC, Plumsteadville, PA, USA) diluted with a zero-air
219 generator (7000 Zero Air Generator, Environics, Tolland, CT, USA). Gases were mixed in a
220 Liquid Calibration Unit (Ionicon Analytik GmbH, Innsbruck, Austria) and water vapor was
221 introduced in the sample to derive the dependence of instrument sensitivity on changing
222 humidity. Formic acid and acetic acid were calibrated before the campaign using liquid standards
223 evaporated and diluted with zero-air in the same Liquid Calibration Unit. Humidity dependence
224 of instrument sensitivity was considered as well. Uncertainty for these species is estimated at
225 40%, with the major source of error being the humidity dependence.

226 Sensitivity for maleic anhydride was estimated using the method by Sekimoto et al. (2017)
227 from its molecular dipole moment and polarizability. The procedure for the calculation was
228 further refined in our previous work (Permar et al., 2021). The uncertainty for this species is
229 estimated to be 50%.

230 Mass spectra were first analyzed with Ionicon's PTR-Viewer software (Version 3.4, Ionicon
231 Analytik). One-minute mass calibrations performed during the campaign were refined using 4
232 ion peaks: m/z 29.997 [NO^+], 59.049 [$\text{C}_3\text{H}_6\text{OH}^+$], 203.943 [$\text{C}_6\text{H}_4\text{IH}^+$], and 330.848 [$\text{C}_6\text{H}_4\text{I}_2\text{H}^+$].
233 Ion counts for each peak in the list were calculated by the PTR-Viewer software through a
234 baseline correction as well as a correction for mass discrimination in the time-of-flight mass
235 spectrometer. The calculated ion counts were then exported for further processing in R.
236 Instrument background was linearly interpolated and subtracted from the data. Each ion was then
237 normalized to the primary ion [H_3O^+] and water cluster ion [$(\text{H}_2\text{O})\text{H}_3\text{O}^+$]. Normalized counts
238 were converted to mixing ratios using the calibrations or theoretical sensitivities described
239 above.

240 2.3 Additional Measurements

241 In addition to the standard UDAQ measurements of CO at UTC (Teledyne API T300U, San
242 Diego, CA, USA), we made concurrent CO measurements using gas chromatography (GC) with
243 a reducing compound photometer (Peak Performer 1 RCP; Peak Laboratories LLC., Mountain
244 View, CA, USA). In this technique, CO eluting from the GC column passes directly into a heated
245 mercuric oxide bed, resulting in liberated mercury vapor, which is subsequently measured via
246 UV light absorption in the photometer cell. Compressed ultra-high purity (UHP) air was used as
247 the carrier gas. Multi-point calibrations are carried out before and after the campaign by dilution
248 of a ppm-level standard (Scott Specialty Gases, USA; stated accuracy $\pm 2\%$) into UHP air. The
249 detection limit for CO is 300 ppt and the time resolution of the data was 3 minutes. The
250 comparisons of two CO measurements, SAMOZA CO and UDAQ CO, show a high correlation
251 during the campaign ($r^2 = 0.86$ and slope = 1.15 with hourly data), with SAMOZA data being
252 systematically higher most of the time. This is likely due to the instrument background drift issue
253 of the regulatory UDAQ CO measurements. It has little policy implication given that SLC CO
254 concentrations rarely exceed the national ambient standard, but it does affect the following
255 analysis. Thus for the rest of the work, we use the CO measurements made by the GC RCP based
256 instrument.

257 Carbonyl samples were collected three times per day using 2,4-dinitrophenylhydrazine
258 (DNPH) cartridges using automatic sampling trays. Cartridges were sampled for three hours with
259 an air flow around 1 L min^{-1} . High-performance liquid chromatography (HPLC) was used to

260 analyze collected samples and quantify the concentrations of 13 different carbonyl species.
261 Samples were kept refrigerated after collection. They were eluted within 14 days of measurement
262 and further analyzed within 30 days. Calibrations were performed with a commercial standard
263 mixture (M-1004; AccuStandard) each day samples were analyzed. Of those carbonyl species
264 measured, 5 overlapped with measurements from the PTR-ToF-MS. In these cases, the PTR-
265 ToF-MS data was prioritized due to the higher temporal resolution. When data from the two
266 methods was compared, the PTR-ToF-MS showed lower concentrations of acetone than DNPH-
267 HPLC, but higher concentrations of formaldehyde, acetaldehyde, and methyl ethyl ketone (Jaffe
268 et al., 2024). The rest of the 8 carbonyl species (all aldehydes) were added to the following
269 analysis and are listed in the supplement (Table S1). The species were used in statistical analysis
270 of VOC composition as well as OH reactivity.

271 Hourly measurements of PM_{2.5}, ozone, and NO_x provided by UDAQ were also used in our
272 analysis. Ozone was measured with a Teledyne T400 instrument. Nitrogen dioxide (NO₂) and
273 nitric oxide (NO) were measured using gas phase chemiluminescence (Teledyne API T200U,
274 San Diego, CA, USA). PM_{2.5} was measured with two Synchronized Hybrid Ambient Real-Time
275 Particulate Monitors (Thermo Environmental 5030i and Thermo Environmental 2025i, Waltham,
276 MA, USA).

277 2.4 Determination of Smoke-influenced Days

278 Each day during the campaign was classified as either being smoke-influenced or smoke-free
279 based on two criteria. The first was the presence of overhead smoke as determined by the
280 National Oceanic and Atmospheric Administration Hazard Mapping System Fire and Smoke
281 Product (<https://www.ospo.noaa.gov/Products/land/hms.html>). It was not used as the only
282 criterion for smoke-influenced days because the presence of overhead smoke does not
283 necessarily indicate that smoke is present at the surface. A second criterion for the concentration
284 of PM_{2.5} at the site was added to ensure the smoke was present at the surface. The mean (6.23 μg
285 m⁻³) and standard deviation (1.85 μg m⁻³) were found for days without any indication of
286 overhead smoke. Days that showed overhead smoke and had a PM_{2.5} concentration one standard
287 deviation above the non-smoke influenced mean were classified as being smoke-influenced.
288 Using these criteria, 10 days out of the campaign were described as smoke-influenced, while 51
289 days were described as smoke-free. As shown in Section 3.2, our VOC analysis still shows some
290 smoke influence in those 51 days. More details on the smoke day identification during
291 SAMOZA can be found at Jaffe et al. (2024).

292 2.5 Positive Matrix Factorization Analysis

293 We performed a positive matrix factorization (PMF) analysis using 26 of the 27 VOCs
294 measured with PTR-ToF-MS for detailed source attribution on the 51 days without the influence
295 of smoke. Dimethyl sulfide was excluded from the analysis due to the majority of the data being
296 below the detection limit. The results were then compared with the traditional approaches
297 including known source tracers and tracer:tracer ratios for qualitative verifications. PMF analysis
298 does not allow for gaps in the data, thus any time that did not have measurements for all 26
299 species was excluded from the model. The analysis is based on the following mathematical
300 model of the data. Let C_{tc} represent the concentration of compound c at time t. We assume there
301 are a number of sources or factors, each one possessing its own temporally-uniform
302 concentration signature. Let X_{sc} represent the mole fraction of compound c in source s. As a
303 mole fraction, X_{sc} is unitless and its sum over compounds is 1: $\sum_c X_{sc} = 1$. We assume that T_{ts}

304 concentration units of source s are present at time t , so that $T_{ts}X_{sc}$ represents the concentration of
 305 compound c derived from source s at time t . Summing over all sources yields the modeled
 306 concentration of compound c at time t :

$$C'_{tc} = \sum_s T_{ts}X_{sc} \quad \#(1)$$

307 The prime is used to distinguish modeled and measured concentrations. In our work, T_{ts} , C_{tc} , and
 308 C'_{tc} are in ppb units. Note that this equation represents a matrix product. Therefore, the
 309 procedure consists of determining two initially unknown matrices T_{ts} and X_{sc} that optimize the fit
 310 between C_{tc} and C'_{tc} (Paatero & Tapper, 1994). Because negative values of T_{ts} and X_{sc} are
 311 physically excluded, the optimization must be carried out subject to the constraints $T_{ts} \geq 0$ and
 312 $X_{sc} \geq 0$.

313 We used the EPA Positive Matrix Factorization 5.0 tool (Norris et al., 2014) to perform
 314 the optimization, with 20 independent runs with unique random seeds to verify that the analysis
 315 converged consistently to the same solution. The number of factors is entered into the model by
 316 the user. One criterion used to determine the appropriate solution was the ratio of Q (goodness-
 317 of-fit parameter calculated with all points) to Q_{expected} (goodness-of-fit parameter excluding
 318 points for which the residual divided by uncertainty is greater than 4), or Q/Q_{expected} .
 319 Increasing the number of factors should decrease Q/Q_{expected} closer to a value of 1 as the error
 320 and/or variability in factor profiles is better accounted for (Ulbrich et al., 2009). The number of
 321 factors was changed between 3 and 9, with a steady decrease in Q/Q_{expected} as the number
 322 increased. However, the solution with 5 factors and Q/Q_{expected} value of 2.56 was chosen for this
 323 analysis due to unreasonable factor splitting as the number increased above 5. This solution
 324 doesn't account for as much error but can be physically interpreted with comparison to
 325 meteorological and concentration data.

326 Uncertainty for each species is given by the user, in the same units as the concentration. This
 327 analysis used uncertainties obtained by multiplying each hourly concentration by its percent
 328 uncertainty stated in section 2.2. The EPA's PMF model includes the option to label a species as
 329 "Strong," "Weak," or "Bad," to further manipulate the uncertainty and therefore how much a
 330 species will contribute to the solution. All species are automatically given the Strong label,
 331 which indicates that the uncertainty supplied by the user will be used. Those species with the
 332 Weak label will have their uncertainties multiplied by 3, so that they have less effect on the
 333 solution the model reaches. Any species with the Bad label will be excluded altogether (Norris et
 334 al. 2014). Before running the model, the signal to noise ratio (S/N) is calculated by the program
 335 for each species. Signal is assigned as 0 for species with greater uncertainty than concentration,
 336 and for all other species the signal is the difference between the concentration and uncertainty.
 337 The S/N is then calculated as the sum of the signal for each data point divided by the number of
 338 data points (Norris et al. 2014).

339 All 26 species showed S/N greater than 0.5 and were therefore left with the default Strong
 340 label. After running the model, the uncertainty-scaled residuals (residual divided by uncertainty)
 341 were evaluated for each species. For those species with residuals outside of the range between -3
 342 and 3, the EPA's standard for a good fit of their model, the observed concentration was
 343 compared with the predicted concentration. Any of those species with large residuals and with an
 344 r^2 value less than 0.6 were labelled as Weak, and the model ran again with the new uncertainties.
 345 From these evaluations, 5 of the 26 included species (furfurals, hexanone, methyl furfurals,
 346 acetonitrile, and formic acid) were labeled as Weak, while the remaining 21 kept the default
 347 label of Strong. Error was estimated using the model's bootstrap method. The model creates a
 348 new dataset by randomly selecting portions of the data until the length of the original dataset is

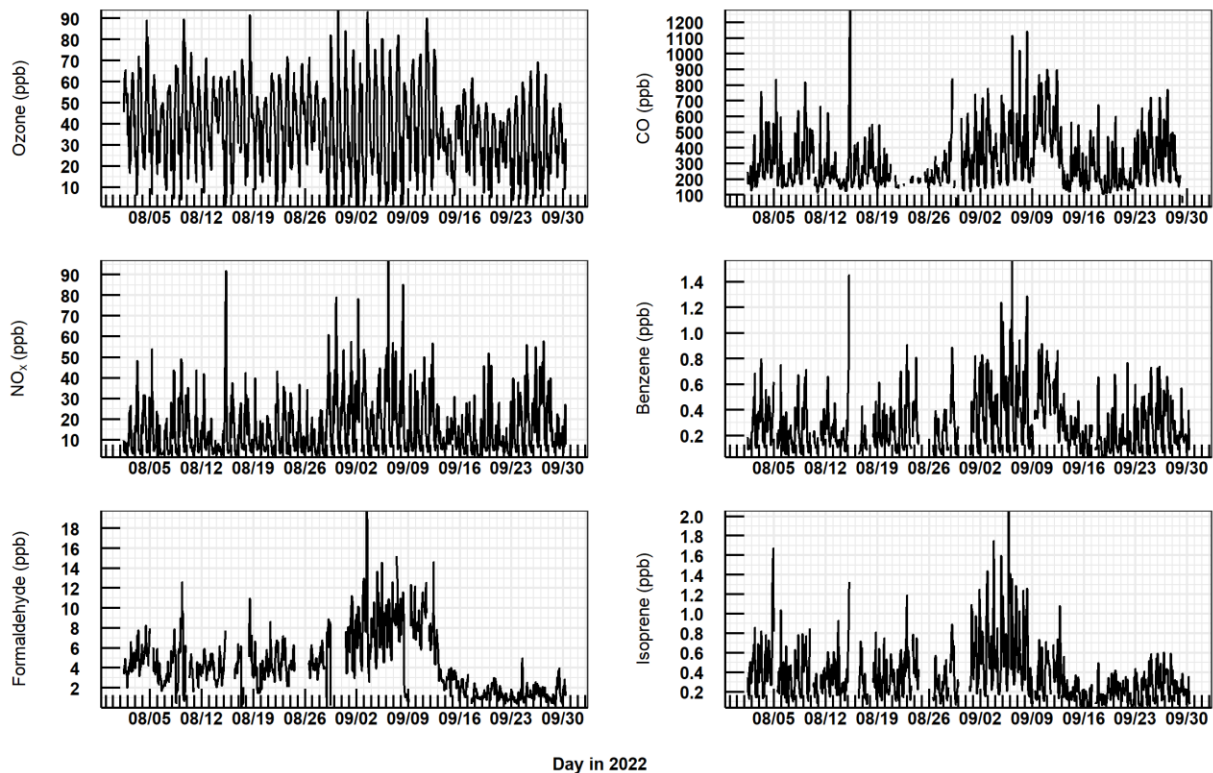
349 reached. The PMF analysis is then performed again with the new dataset, and each new factor is
350 mapped against the factor from the original dataset that it has the highest correlation with, above
351 a threshold defined by the user (Norris et al., 2014). For this analysis, a minimum correlation
352 value of 0.72 yielded no unmapped factors. Factor 2 mapped to its original factor for 52% of the
353 runs, and all other factors mapped to the original more than 90%.

354 **3 Results and Discussion**

355 3.1 Overview: VOC Abundance, Composition, and Reactivity in SLC

356 Figure 1 shows the meteorological data and Figure 2 shows the time series for ozone,
357 CO, NO_x, and selected VOCs measured during SAMOZA. August-September 2022 at the SLC
358 airport was slightly warmer (~2°C) than the decadal mean and was a relatively modest smoke
359 year with 10 out of 61 days identified as smoke-influenced days during SMAOZA (Jaffe et al.,
360 2024). As mentioned above, the average daily air temperature was ~10°C higher in the first ~2/3
361 of the SAMOZA period than the last 1/3 of days, reflecting the hot/dry summer to cool/wet
362 autumn seasonal transition in SLC starting around September 9, 2022. Interestingly, five out of
363 the 10 smoke days occurred in September when the air temperature was cooler, including the
364 four of the smokiest days (9th-12th September 2022) that reached the highest daily mean PM_{2.5} of
365 25.3-33.8 ug/m³ during SAMOZA (Jaffe et al., 2024). It is likely that those smoke days were
366 influenced by large fires burning in other western states and transported to SLC. Our VOC
367 analysis shown later also suggests the smoke was indeed aged during SAMOZA (Section 3.3).
368 Ozone concentrations averaged 37 ± 19 ppb (average \pm standard deviation), with hourly peaks
369 above 70 ppb on 22 days, and five MDA8 exceedance days (8/4, 8/9, 9/3, 9/7, 9/11). The
370 corresponding averaged NO_x was 16 ± 13 ppb at the UDAQ UTC site, comparable to the level in
371 other non-attainment urban areas in the western US (i.e., May-September averaged ~15 ppb
372 across 20 west urban sites in Jaffe et al. 2022).

373 Over the two months of SAMOZA, the total measured VOC mixing ratio averages $36 \pm$
374 24 ppb (or 82 ± 55 ppbC). Wind rose plots in Figure S1 show the wind direction separated by
375 day and night. These plots suggest that the highest frequency of wind came from the southeastern
376 direction, which mostly occurred at night (18:00 – 06:00 Mountain Standard Time). This
377 direction was also correlated with the higher mean concentration of VOCs during the campaign
378 compared to the western wind, at 35 ± 26 ppb. As the UDAQ UTC location was situated in the
379 northwestern part of Salt Lake City, southeasterly winds could be influenced by various urban
380 sources in the city. A significant amount of wind also came from the west and northwestern
381 direction, mostly also occurring during daytime when the photochemical loss of VOCs is strong
382 and the atmospheric boundary layer is high. It resulted in a lower total measured VOC
383 abundance (19.4 ± 13 ppb) under such conditions when the site is likely influenced by the
384 industrial sites and Salt Lake City International Airport.



385
 386 **Figure 2.** Time series plots for hourly concentrations of ozone, CO, NO_x, and selected VOCs
 387 including benzene, formaldehyde, and isoprene between August 1 and September 30, 2022
 388 during the SAMOZA field campaign.

389
 390 The overall VOC composition during SAMOZA appears to have two distinct
 391 distributions likely controlled by air temperature. VOC species mostly driven by primary
 392 anthropogenic emissions, such as benzene, toluene, and ethanol, as well as CO and NO_x, do not
 393 show significant changes when the air temperature dropped in the middle of September (Figure
 394 2). However, those compounds with biogenic origin or photochemical sources exhibit a similar
 395 pattern of decreased concentration by 35-50% starting around September 12th when daily
 396 average air temperature was about 10°C lower than the previous period. Isoprene and
 397 monoterpenes are examples that are often majorly influenced by biogenic sources, though urban
 398 sources or measurement interferences could play a nonnegligible role (Borbon et al., 2023;
 399 Coggon et al., 2023; Hu et al., 2015b). Formaldehyde and other species like methyl vinyl ketone
 400 and methacrolein (MVK + MACR) with large photochemical productions also show this pattern
 401 of decreased concentration with temperature towards the end of the campaign. Most of the
 402 mixing ratios for these species (formaldehyde, methyl ethyl ketone [MEK], MVK + MACR, and
 403 acetaldehyde) were 35-45% lower compared to the average of the first 2/3 of the period. When
 404 the daily average mixing ratio for the VOCs was correlated against the daily average
 405 temperature, those secondary and biogenic species showed significantly higher r^2 values (e.g.,
 406 0.58 for formaldehyde, 0.64 for isoprene) than the species that did not show decreased
 407 temperature after September 9th (e.g., 0.18 for benzene, 0.13 for ethanol), further indicating that
 408 temperature played a role in controlling VOCs from biogenic sources or secondary production.

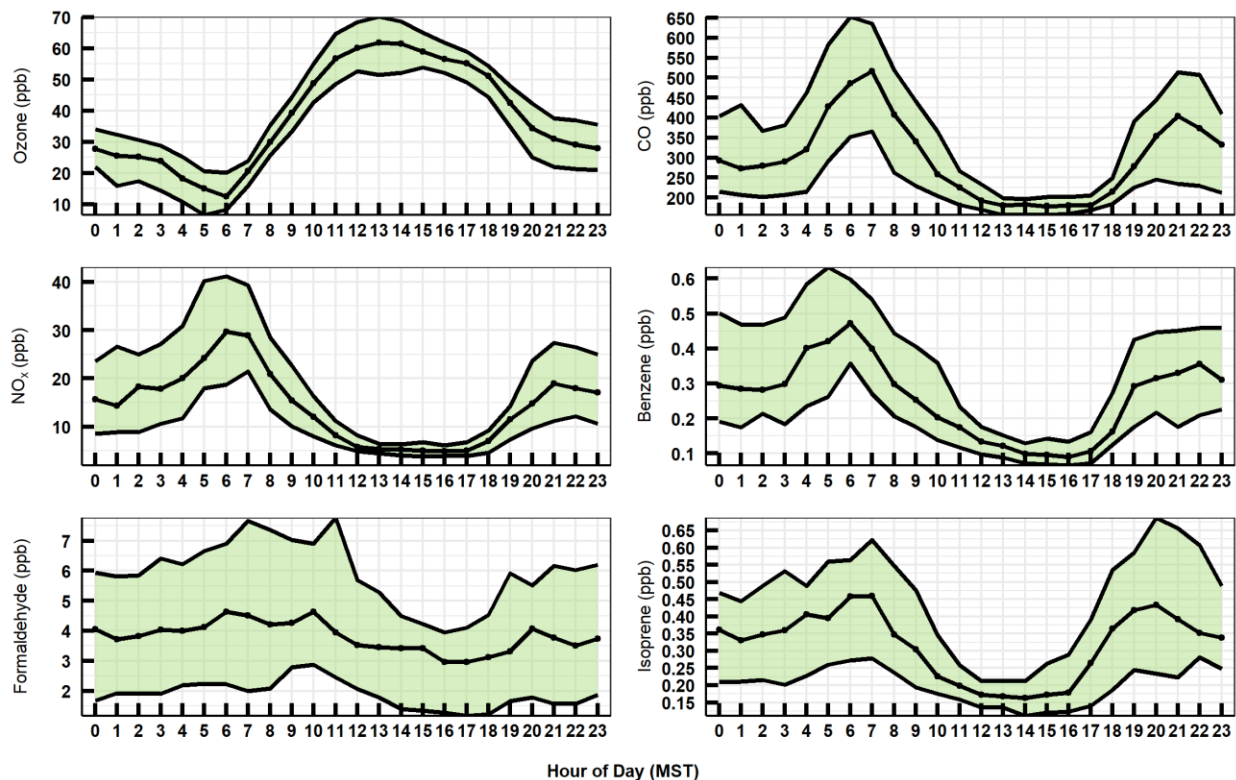
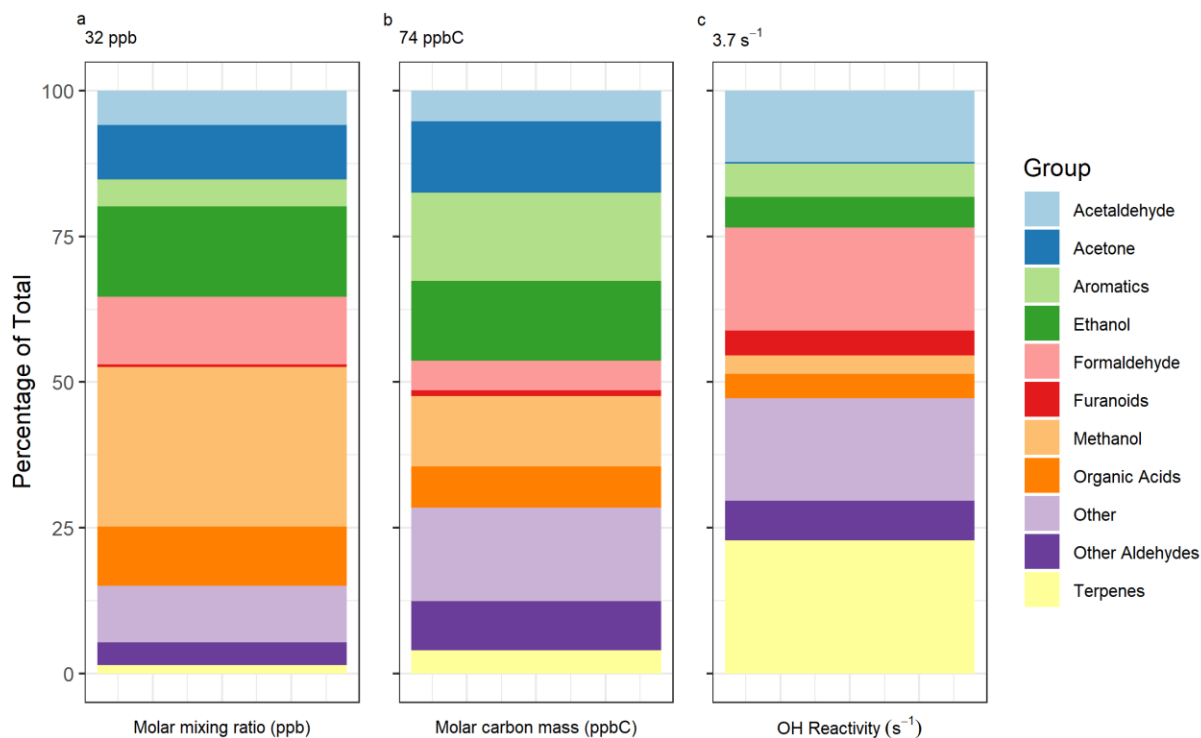


Figure 3. Diurnal variations of ozone, CO, NO_x, benzene, formaldehyde, and isoprene observed during SAMOZA (August 1 to September 30, 2022). The thick middle lines indicate the median mixing ratio, with the shaded areas as 25th-75th percentiles of the observations.

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The diurnal patterns for primary anthropogenic VOCs mostly showed a similar pattern, with a large peak in concentration early in the morning (06:00 – 09:00 Mountain Standard Time), rapid decreasing during the day, and then increasing at night (Figure 3). Such diurnal profiles have been widely documented in urban sites (e.g. Bryant et al., 2023; Coggon et al., 2018; de Gouw et al., 2017; Li et al., 2022), reflecting the combined effects of emission strength and atmospheric boundary layer dynamics, while the ratios of nighttime and daytime values are often determined by their lifetimes (de Gouw et al. 2017). The group of C₆-C₁₀ aromatic compounds is one such example. These compounds are highly correlated to each other (R^2 ranging from 0.75 to 0.94) and show the same diurnal patterns but have averaged nighttime: daytime ratios (22:00-06:00 / 10:00-18:00) increasing from 2.2 (benzene) to 4.3 (C₁₀ aromatics) following the increasing order in their OH rate constants. Oxygenated VOCs, such as formaldehyde, formic acid, and MVK+MACR, had a slight peak in the morning but did not show the same large decrease during the day as in other species. This is due to their sources, either biogenic or photochemical, being enhanced during the day when the solar radiation is strong and temperature is high, even though they may also have primary anthropogenic emissions in urban areas. Isoprene showed a similar diurnal pattern to benzene and CO, which could indicate a strong anthropogenic influence in SLC in addition to biogenic emissions. Later in Section 3.2, we will present a more quantitative analysis of their sources in SLC.



432
 433 **Figure 4.** The campaign-averaged contribution of major VOC species/groups to total measured
 434 VOCs in SLC during SAMOZA as a) molar mixing ratio in ppb, b) molar carbon mass in ppbC,
 435 and c) daytime OH reactivity in s^{-1} . For the VOC groups, Aromatics include benzene, toluene,
 436 and C₈, C₉ and C₁₀ aromatics; Furanoids include furan, methylfuran, furfural, and methyl
 437 furfural; Organic Acids include formic acid and acetic acid; Other includes propyne, acetonitrile,
 438 butene, dimethyl sulfide, MVK + MACR, methyl ethyl ketone (MEK), maleic anhydride,
 439 hexanone, and D5 siloxane; Other Aldehydes includes the aldehydes measured with DNP
 440 cartridge: acrolein, propionaldehyde, crotonaldehyde, n-butyraldehyde, valeraldehyde,
 441 hexaldehyde, benzaldehyde, and m-tolualdehyde; Terpenes include isoprene and monoterpenes.
 442 The average values for each metric are shown at the top of the bars. VOC data from 35
 443 species/masses are from both PTR-ToF-MS and DNP/HPLC measurements. See Table S1 for
 444 details.

445
 446 Figure 4 shows the percent contribution of VOCs for the SAMOZA campaign average,
 447 separated into major functional groups and major individual contributors, to overall molar mass
 448 (ppb), molar carbon mass (ppbC), and OH reactivity (s^{-1}). Here we integrated the DNP/HPLC
 449 measurements for a more complete list of species; all 8 species are included under the Other
 450 Aldehydes category. The total measured VOCs from 35 species/masses averaged 32 ± 24 ppb,
 451 with a median of 27 ppb (Interquartile range (IQR) of 27 ppb, and max 141 ppb). This
 452 corresponds to 74 ± 57 ppbC (median 58 ppbC, IQR 64 ppbC, and max 345 ppbC).

453 OH reactivity was calculated for VOCs for Figure 4 as well as for NO_x for later analyses
 454 according to the equation

$$OHR = k_{OH+X}[X]\#(2)$$

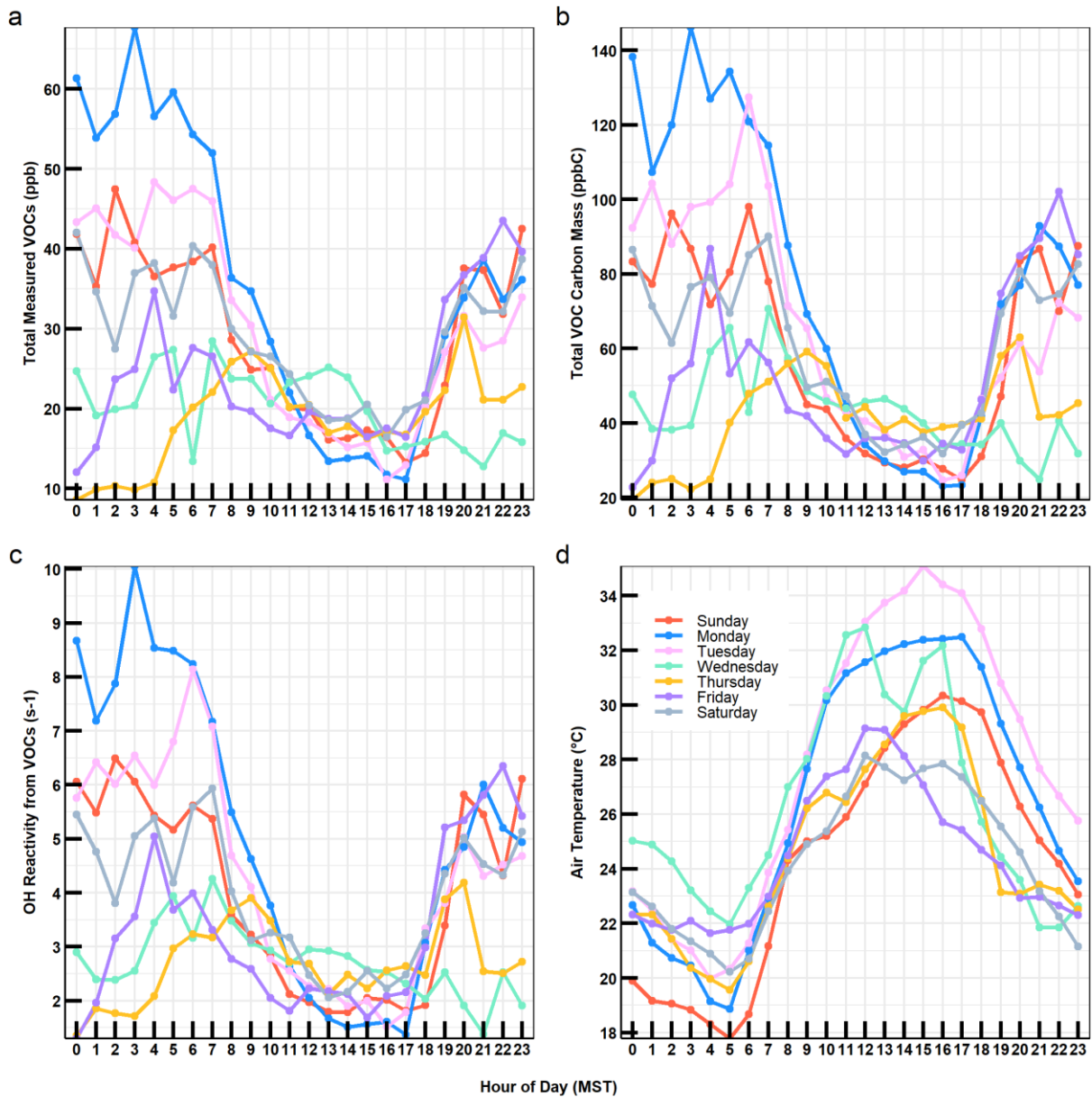
455 Where k_{OH+X} is the rate constant for the reaction of species X with the OH radical and [X] is the
 456 measured species concentration. Rate constants were taken from Permar et al. 2023, with

457 additional rate constants not included (propyne, D5 siloxane, and NO_x) found using the NIST
458 chemical kinetics database (<https://kinetics.nist.gov/kinetics/>). Concentrations for each species
459 were converted to number density in molecules per cm³ using the field-measured temperature
460 and pressure. The total calculated OHR from measured VOCs averaged $3.7 \pm 3.0 \text{ s}^{-1}$, with a
461 median of 2.9 s^{-1} (IQR 2.7 s^{-1} and max 20.7 s^{-1}) for the SAMOZA campaign.

462 Methanol was the most abundant VOC in SLC in terms of mixing ratio (27% of total
463 measured VOCs), followed by ethanol (16%). This is consistent with other urban areas such as in
464 India (Kalbande et al., 2022) and France (Simon et al., 2023). In contrast, ethanol was found to
465 be the most abundant VOC in Guangzhou, China, followed by methanol (Li et al., 2022).
466 Methanol is thought to have a large biogenic origin even in urban atmospheres (de Gouw et al.,
467 2005; Hu et al., 2011; MacDonald & Fall 1993; Salisbury et al., 2003). Although ethanol can be
468 found in fuel, it has been shown to have significant biogenic sources in the western US (Millet et
469 al., 2012) and has recently been attributed to VCPs in urban areas (McDonald et al., 2018). Also
470 among the most abundant VOCs were acetone (9%), acetic acid and acetaldehyde (both 6%), and
471 formaldehyde (12%).

472 When weighted by the number of carbons, methanol is not as important a contributor,
473 only accounting for 12% of the total measured VOC carbon. Aromatic compounds saw a large
474 increase in contribution by this metric (5% of total measured VOCs to 15% of total measured
475 VOC carbon). VOCs in the Other category, such as acetonitrile, butenes, and propyne, also saw
476 an increase in contribution in this way (10% total measured VOCs compared to 16% total
477 measured VOC carbon).

478 In terms of OH reactivity, isoprene and monoterpenes, lumped as Terpenes in Figure 4,
479 were the most abundant (23% of total VOC OH reactivity) despite making up less than 2% of the
480 total mixing ratio due to their rapid reactions with OH radicals. Isoprene has been shown to be a
481 large contributor to VOC OH reactivity in urban regions of South Korea (Sanchez et al., 2021)
482 and Canada (Stroud et al., 2008). In other urban areas, OH reactivity from VOCs has been more
483 largely influenced by aromatics and alkenes from traffic and industrial sources (Gilman et al.,
484 2009) or OVOCs such as carbonyls and alcohols (Hansen et al., 2021). Similarly to the biogenic
485 influence on VOC abundance, Salt Lake City appears to have a large influence from biogenic
486 emissions on daytime photochemistry. Formaldehyde and acetaldehyde were also important
487 contributors to OH reactivity at 18% and 12% of the total, respectively. Furanoids including
488 furan, methyl furans, furfural, and methylfurfurals, were some of the least abundant VOCs by all
489 three metrics (i.e., <1% of total measured VOCs), reflecting their relatively small urban and BB
490 emissions during SAMOZA.

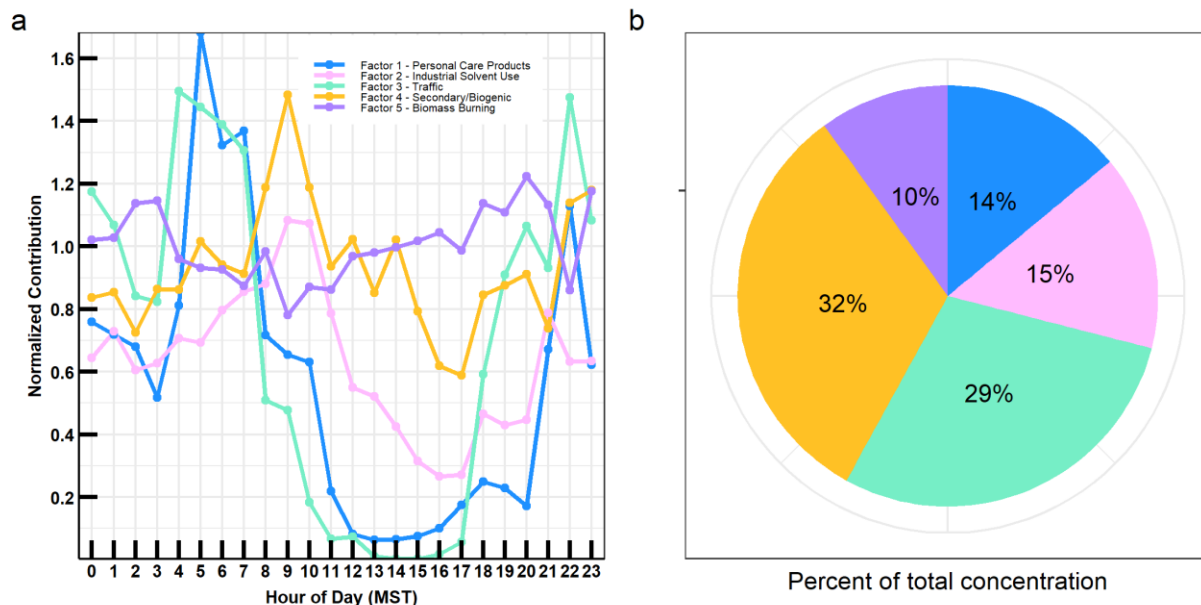


491
 492 **Figure 5.** Diurnal patterns for a) the median total measured VOC concentration in ppb, b) the
 493 median total ppbC, c) the median total OHR from VOCs in s⁻¹, and d) the median air temperature
 494 averaged by day of the week for the 51 smoke-free days over the two-month SAMOZA period.
 495 This includes 7 Sundays, 8 Mondays, 9 Tuesdays, 7 Wednesdays, 7 Thursdays, 7 Fridays, and 6
 496 Saturdays.

497
 498 Figure 5 shows the diurnal pattern for total median VOC concentrations and OH
 499 reactivity separated by the days of the week. VOC abundances peak in the evening and early
 500 morning and show a large dip in the middle of the day as temperatures rise and the atmospheric
 501 boundary layer breaks down. The highest daily concentration of VOCs is present on Mondays
 502 and Tuesdays (35 ± 27 ppb vs. 31 ± 24 ppb). Thursdays show the lowest daily VOC
 503 concentration (21 ± 18 ppb), with the lowest concentrations observed during the night, not
 504 consistent with the pattern of the other days of the week, when the lowest concentrations were

505 observed between 12:00 and 18:00 MST. The concentration of VOCs is consistent with the
 506 patterns of NO_x also measured during SAMOZA (Jaffe et al., 2024). OHR and VOC carbon mass
 507 showed similar patterns, with Mondays and Tuesdays being the highest ($5.2 \pm 4.2 \text{ s}^{-1}$ vs. $4.4 \pm$
 508 3.8 s^{-1} ; $80 \pm 65 \text{ ppbC}$ vs. $67 \pm 58 \text{ ppbC}$) and Thursdays being the lowest ($3.0 \pm 2.6 \text{ s}^{-1}$; 45 ± 39
 509 ppbC). The exact reasons for the patterns of the days of the week are unknown but it may be
 510 related to emission patterns and meteorological conditions. Mondays and Tuesdays indeed show
 511 higher daytime (06:00 to 18:00 MST) temperatures than the other days, averaging $29 \pm 5.1^\circ\text{C}$
 512 compared to $26 \pm 5.5^\circ\text{C}$. However, when plotted against temperature, the median values did not
 513 show any clear pattern, so the role of temperature is not well known. When we removed those
 514 Mondays and Tuesdays with particularly high values of total measured VOCs to investigate if
 515 the pattern was affected by them, as was the case for NO_x in Jaffe et al. (2024), we found such
 516 omission made little difference in the diurnal patterns. When examining the weekday vs.
 517 weekend difference, we found there are no statistically significant differences in total VOC
 518 abundance or OH reactivity (Section 3.4). Future work to include longer periods of observation
 519 will help improve statistics and examine if there are meaningful emission differences in the days
 520 of the week in SLC.

521 3.2 Sources of Major VOCs



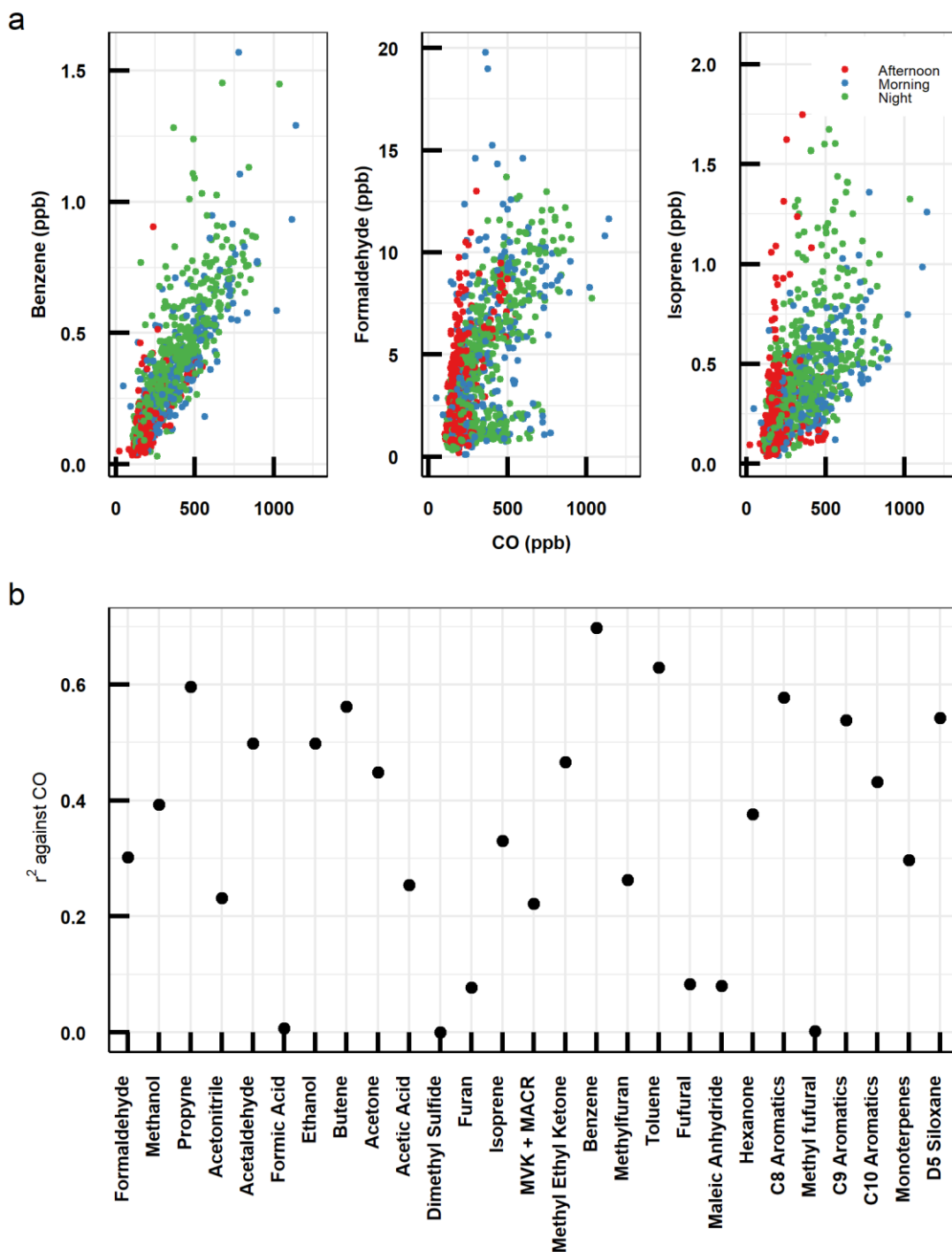
522 **Figure 6.** a) Diurnal pattern of median normalized concentration for each factor in the PMF
 523 analysis: Factor 1 – Personal Care Products dominated, Factor 2 – Industrial Solvent Use, Factor
 524 3 – Traffic dominated, Factor 4 – Secondary/Biogenic, Factor 5 – Biomass Burning. b) Pie chart
 525 showing the percentage of total concentration contributed by each factor.
 526

527

528 Here we investigate the main sources of VOCs in SLC using PMF analysis. The identity
529 of the factors was assigned based on features in the time series and correlation with other tracers
530 and meteorological conditions such as temperature and solar radiation. The diurnal patterns of
531 the factors are given in Figure 6a and were also used in identification. The abundance of species
532 within the factor profiles (shown in Figure S2) was used as the final confirmation of the source.
533 Factors 1 and 3 are dominated by Personal Care Products and Traffic sources, respectively. They
534 correlate well with NO_x and CO and have a similar diurnal pattern, with a large dip in the day in
535 the diurnal pattern, discussed previously as an indicator of anthropogenic influence. Notably,
536 there are significant amounts of aromatic compounds attributed to the personal care product
537 factor, and significant amounts of D5 siloxane and monoterpenes attributed to the Traffic factor.
538 Co-emission of these compounds during peak traffic times may explain the difficulty in fully
539 extricating these two sources in PMF (Coggon et al., 2018). Ultimately, Factor 1 was assigned as
540 Personal Care Products dominated due to the high abundance of monoterpenes and low
541 abundance of benzene, and Factor 3 was assigned as Traffic dominated due to high amounts of
542 both benzene and toluene.

543 Factor 2 is assigned as Industrial Solvent Use due to a similar diurnal pattern to the other
544 anthropogenic factors, albeit with a shallower dip during the day in the diurnal pattern as well as
545 high levels of OVOCs like MEK and formaldehyde. Factor 4 includes species produced in
546 photochemical reactions and emitted by biogenic sources. The diurnal pattern shows no large
547 decrease in concentration during daylight hours likely due to increased chemical reactions and
548 plant activity. This factor also showed a large decrease after September 9 when the air
549 temperature dropped 10°C, reflecting the seasonal reduction in photochemical and plant
550 biological activities. Important species for this factor include formaldehyde, isoprene, and its
551 oxidation products MVK+MACR. Factor 5 is labeled as Biomass Burning due to containing
552 much of the furanoid species and acetonitrile and because it does not have a distinct diurnal
553 pattern, with the median concentration staying similar throughout the day.

554 The pie chart in Figure 6b shows large contributions from Secondary/Biogenic sources
555 (~32%). According to this model, traffic emissions are still a large contributor to VOCs in SLC
556 at 29% of total abundance. Using only Factors 1, 2, and 3 as primary urban emissions, traffic
557 contributes 50% of the anthropogenic VOCs, while Personal Care Products and Industrial
558 Solvent Use are roughly equal at 25%. Since the latter two factors both fall under the category of
559 VCPs, this means that traffic and VCPs contribute about the same to anthropogenic VOCs in
560 SLC, similar to what was estimated in the 2020 NEI. We note that PMF results shown here,
561 though somewhat quantitative, could be subject to other compounding factors, such as the
562 Personal Care Products and Traffic factors are not able to be fully resolved due to their similar
563 emission patterns.



564

565 **Figure 7.** a) Benzene, formaldehyde, and isoprene mixing ratios plotted against observed CO, all
 566 in ppb. The data points are colored by the time of day: morning was defined as 06:00-12:00
 567 MST, afternoon was defined as 12:00-18:00 MST, and night was defined as 18:00-06:00 MST.
 568 b) The r^2 value for the reduced major axis regression of each VOC against CO measured during
 569 SAMOZA.

570

571 We further use source tracer analysis to test if it can qualitatively confirm the PMF
572 findings or not, with a particular focus on well-known emission tracers or signatures. Figure 7
573 shows benzene, formaldehyde, and isoprene correlated with carbon monoxide (CO), a vehicle
574 tracer that is not used in PMF, and colored by time of day. Correlation coefficient values are also
575 given for all 27 VOC species measured by PTR-ToF-MS. Benzene and other aromatic
576 compounds were seen to correlate best with CO (0.70 and 0.63, respectively), with no apparent
577 difference in population based on time of day. Although they can have several anthropogenic
578 sources, aromatic compounds are present in vehicle fuel and known to be co-emitted with CO.
579 Formaldehyde and other compounds with secondary production sources are seen to have two
580 populations and lower correlation with CO, with the highest slopes seen in the afternoon when
581 chemical production is highest. Within the PMF model, 54% of formaldehyde was assigned to
582 the Secondary/Biogenic factor, while the rest was split between the three anthropogenic factors.
583 Isoprene and other biogenic VOCs show a similar low correlation with CO and multiple
584 populations. Slopes were not as clearly separated for these compounds as for formaldehyde. A
585 large amount of isoprene comes from biogenic sources as stated previously, but there are some
586 anthropogenic sources as well (Bryant et al., 2023; Khan et al., 2018; Wagner & Kuttler, 2014).
587 From the PMF analysis, 40% of isoprene was assigned to the Secondary/Biogenic factor, and
588 50% was mapped to the three anthropogenic factors. When plotted against CO, the Traffic factor
589 showed the highest correlation ($r^2 = 0.58$), while the other anthropogenic factors had lower
590 correlation (0.23 for Personal Care Products and 0.19 for Industrial Solvent Use), and the natural
591 factors had the lowest correlation (<0.01 for Secondary/Biogenic and Biomass Burning). The
592 considerable number of VOCs correlating with CO corroborates the PMF findings that Traffic is
593 an important VOC source in SLC.

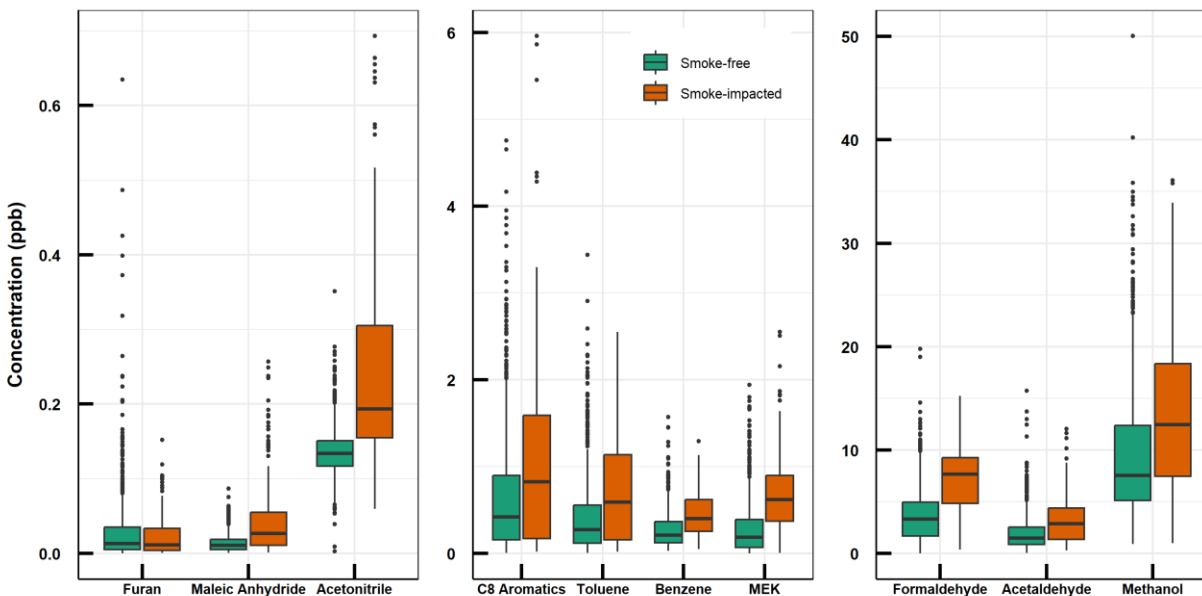
594 The ratio of toluene to benzene is another important emission signature for traffic,
595 biomass burning, and industrial sources. Figure S3 shows the correlation between the two
596 species. The ratio of toluene to benzene in SLC during SAMOZA was 2.02. Ratios higher than 1
597 are commonly seen in urban settings with significant solvent use in the area, and ratios between
598 1-10 are commonly seen when traffic is the major VOC source (Zhang et al., 2016). Data here is
599 given for the full two-month campaign. Separating the data by smoke-influenced and smoke-free
600 days did not change the ratio significantly, with values of 2.09 and 2.06, respectively. This ratio
601 also corroborates the assignment of the anthropogenic factors in the PMF model. The Traffic
602 factor had a lower toluene/benzene ratio (1.55) than that of Personal Care Products (2.44) or
603 Solvent Use (2.43).

604 In addition, isoprene and monoterpenes are known to be emitted from biological
605 processes in plants. Figure S4 shows the correlations of each compound with temperature.
606 Isoprene concentration has a strong increase with temperature, while monoterpenes reach the
607 highest concentration at 22°C with a decrease at higher and lower temperatures. When correlated
608 with CO, colored by temperature, isoprene shows some separation in correlation. Higher
609 temperatures led to a higher slope. Monoterpenes, in contrast, do not show this same separation.
610 This agrees with the PMF analysis, wherein roughly 76% of the concentration of monoterpenes
611 was attributed to the factors for Personal Care Products and Traffic.

612 D5 Siloxane can be used as a tracer for personal care products (Gkatzelis et al., 2021),
613 but shows similar emission patterns as traffic. In Salt Lake City, we see a strong correlation
614 between D5 and benzene during SAMOZA ($r^2 = 0.72$; Figure S5). They show similar diurnal
615 patterns as well, with peaks early in the morning as traffic begins. The two are not perfectly

616 correlated, as humans spend much of their time indoors rather than driving, but there is some
 617 overlap there, similar to the emission features derived by Coggon et al. (2018). As discussed
 618 above, PMF cannot tell such co-emitted features from two different sources, with about 60% of
 619 D5 siloxane and 40% of benzene mapped to Traffic, and 20% of each to the Personal Care
 620 Products dominated factor. Therefore the PMF results here should be considered semi-
 621 quantitative. Future studies including more source tracers or source measurements from mobile
 622 platforms can help better understand the the complexity of urban emissions in SLC.

623 3.3 Impact of Wildfire Smoke



624 **Figure 8.** Box and whisker plot showing the concentration of key wildfire-related species (furan,
 625 maleic anhydride, and acetonitrile) as well as Hazardous Air Pollutants (HAPs) identified by the
 626 EPA (C8 aromatics, toluene, benzene, MEK, acetaldehyde, formaldehyde, and methanol) in ppb.
 627 The concentration is shown for the 51 smoke-free days and the 10 smoke-influenced days during
 628 SAMOZA.
 629

630

631 Figure 8 shows key species on smoke-influenced days and smoke-free days during
 632 SAMOZA. The average abundance of total VOCs increased from 33 ± 21 to 50 ± 28 ppb when
 633 fire smoke impacted the area. Most individual VOCs (26 out of 35) also show enhancements
 634 during smoke-influenced days, with the largest enhancement being present for maleic anhydride
 635 and MEK at over 100% increases compared to smoke-free days. Acetonitrile is enhanced by
 636 85% in the presence of smoke. Acetonitrile has been widely used in past studies as a biomass
 637 burning tracer, even though it also has a smaller anthropogenic source that is also reflected in our
 638 PMF analysis (Huangfu et al., 2021). Furan shows slightly lower concentration on days that are
 639 smoke-enhanced, while maleic anhydride shows an enhancement in smoke of 400%. Furan is
 640 emitted directly from biomass burning and has a short lifetime (estimated 7 hours assuming OH
 641 concentrations of 1.0×10^6 molecules/ cm^3). It reacts rapidly with OH radicals to form oxidation
 642 products including maleic anhydride, which has a longer lifetime (estimated 8 days) (Bierbach et

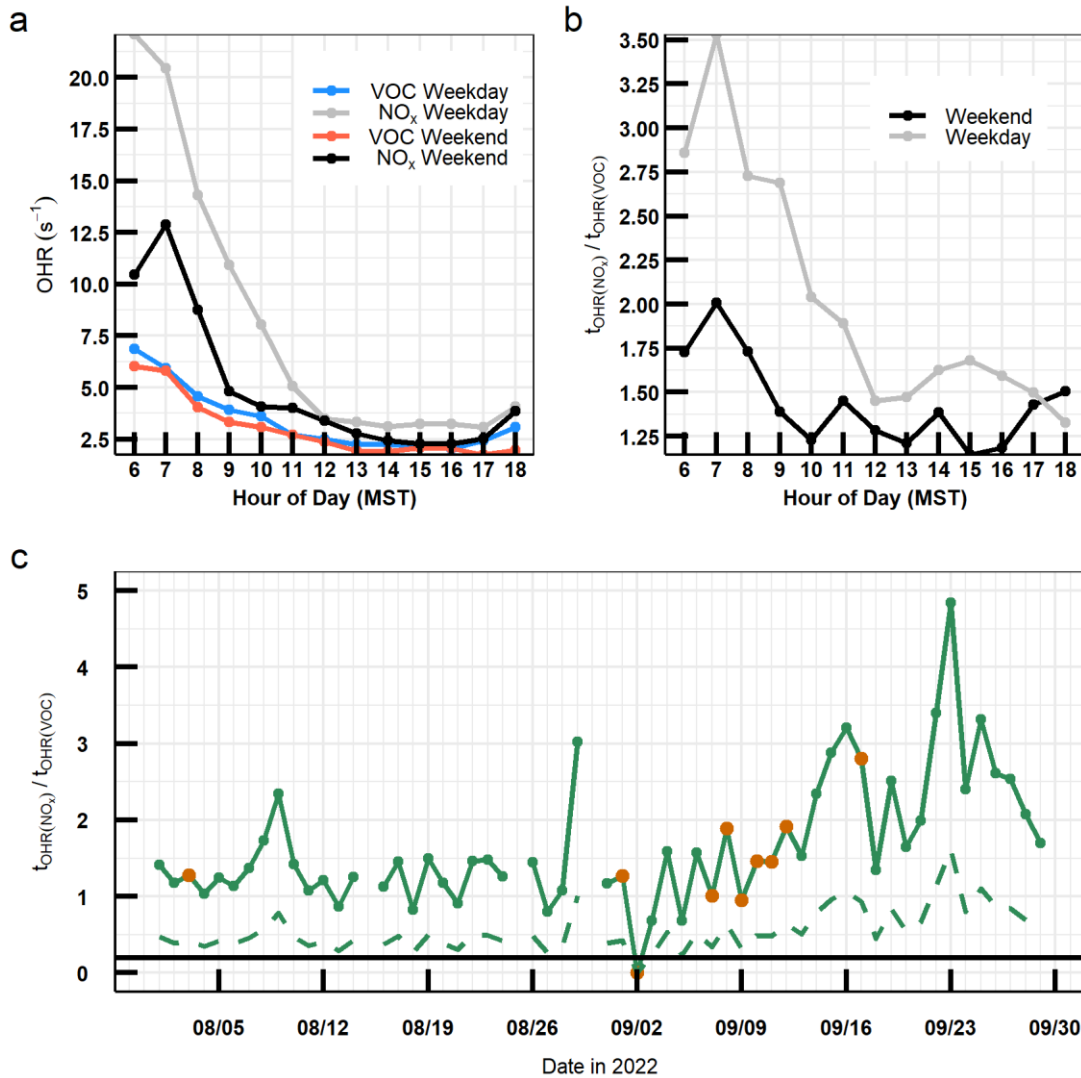
643 al., 1994, 1995). The high maleic anhydride and low furan levels indicate that the smoke in Salt
644 Lake City during SAMOZA had photochemically aged before reaching the urban area, thus
645 originating from regional smoke rather than a local fire emission. Eight species included in this
646 analysis (acetaldehyde, acetonitrile, benzene, C₈ aromatics, maleic anhydride, methanol, toluene,
647 and formaldehyde) are on the EPA's list of HAPs, all shown in Figure 8. Their abundances were
648 increased during smoke-influenced days by 45-217%. Those species had an average
649 enhancement of 88% when smoke was present in Salt Lake City during SAMOZA. CO and
650 PM_{2.5}, criteria air pollutants regulated by the EPA, also showed enhancement in smoke (88% and
651 174%, respectively), while ozone had a much lower increase at only 5%. Mixing ratios for all
652 observed species during smoke-free and smoke-influenced days are found in Table S1. Solar
653 radiation and temperature were additionally analyzed between the two groups, but because there
654 was no statistically significant difference for either measurement ($1.58 \times 10^5 \pm 1.97 \times 10^5 \text{ W/m}^2$
655 vs. $1.71 \times 10^5 \pm 2.07 \times 10^5 \text{ W/m}^2$; $25 \pm 5.6^\circ\text{C}$ vs. $25 \pm 6.3^\circ\text{C}$), they are not thought to contribute
656 significantly to the changes in VOC concentration.

657 Figure S6 shows the total OH reactivity from VOCs increased by 50% during smoke-
658 influenced days compared to smoke-free days (5.1 vs 3.4 s^{-1}). Terpenes including isoprene and
659 monoterpenes that come largely from biogenic sources, are not as important during smoke-
660 influenced periods due to formaldehyde and other oxygenated VOCs becoming more important
661 contributors to OH reactivity during these days. Furanoid OH reactivity remained similar in both
662 cases, which caused the contribution to total OH reactivity to decrease as aged fire smoke
663 traveled to SLC during SAMOZA. Other VOC groups also show similar reactivity in both
664 periods in spite of increased concentrations.

665 The impact of smoke on ozone production during SAMOZA was investigated in a
666 separate publication, and thus is not a focus in this study. In brief, Ninneman et al. (2023) used
667 observationally constrained photochemical box modeling for four case study days: including two
668 smoke-free weekday and weekend days, and two smoke-influenced weekday and weekend days
669 (Ninneman et al., 2023). We found that predicted levels of ozone were similar between the
670 smoke-free and smoke-influenced days, however the sensitivity of ozone production was seen to
671 be affected significantly by smoke. In both cases, ozone was not sensitive to reductions in NO_x
672 without reductions of 75% or more. The difference was that ozone on smoke-free days was
673 sensitive to VOCs in general, while ozone on smoke-influenced days was sensitive to fire VOCs
674 to a much larger degree than any anthropogenic VOCs. While smoke did not cause a significant
675 increase in ozone concentration during SAMOZA, smoke-emitted VOCs controlled the
676 chemistry of ozone production when they reached the area. We refer to Ninneman et al (2023)
677 for details of smoke impact on ozone, while in the next section, our analysis focuses on ozone
678 photochemistry on smoke-free days which were the majority of the SAMOZA campaign.

679 3.4 Implications for Urban Photochemistry

680 Here we examine the transition of photochemistry in SLC during SAMOZA. We
681 calculate the total OH reactivity (OHR) from NO_x ($t\text{OHR}_{\text{NO}_x}$) and from VOCs ($t\text{OHR}_{\text{VOC}}$), and
682 their ratio $\theta = t\text{OHR}_{\text{NO}_x}/t\text{OHR}_{\text{VOC}}$ from SAMOZA measurements. The ratio θ was proposed as a
683 metric to define the ozone production regime by Kirchner et al. (2001). By this metric, the
684 production regime is VOC-limited when $\theta > 0.2$, NO_x-limited when $\theta < 0.01$, and transitional when
685 θ is between 0.01 and 0.2. This metric has been applied in previous studies in the western US to
686 show that VOCs emitted by smoke changed the ozone production regime in urban areas (Liang
687 et al., 2022; Permar et al., 2023).



688

689 **Figure 9.** Daytime (06:00 – 18:00 MST) variations of OHR in s⁻¹ a) for NO_x and VOCs
 690 differentiated by weekends and weekdays and b) for the ratio of $t_{\text{OHR}}(\text{NO}_x) / t_{\text{OHR}}(\text{VOC})$ or θ
 691 differentiated by weekends and weekdays. Both plots only include the 51 days not influenced by
 692 smoke. c) The daytime (06:00 – 18:00 MST) ratio of $t_{\text{OHR}}(\text{NO}_x) / t_{\text{OHR}}(\text{VOC})$ for each day during
 693 SAMOZA. Smoke-influenced days are highlighted in orange. The black horizontal line indicates
 694 where θ is equal to 0.2, the threshold between the VOC-limited and transitional regimes. The
 695 green dashed line shows the estimated lower bound of θ , by accounting for the potential
 696 influence from unmeasured VOCs; See text for details.

697

698 Figure 9 shows the daytime θ each day for the duration of SAMOZA, with smoke-
 699 influenced days highlighted in orange. θ was larger than 0.2 throughout the SAMOZA, and the
 700 ozone production is therefore estimated to have been VOC-limited. This is consistent with the
 701 findings of our modeling study (Ninneman et al., 2023). When separated by smoke-influenced
 702 and smoke-free days, there was little change in the regime, evident in the time series. On smoke-
 703 free days, θ averaged 1.8 ± 1.1 , while on smoke-influenced days, θ averaged 1.4 ± 0.73 , both

704 well above the threshold for the VOC-limited regime. This contrasts previous studies such as
705 Liang et al. (2022) and Permar et al. (2023), which showed wildfire smoke lowering θ due to
706 enhanced VOCs and shifting the ozone production regime from VOC-limited to a NO_x -limited or
707 transitional regime. In the case of SAMOZA, NO_x concentrations were enhanced along with
708 VOC concentrations during smoke-influenced days, which resulted in an increase in $\text{tOHR}_{\text{NO}_x}$
709 from 4.71 s^{-1} on smoke-free days to 6.72 s^{-1} on smoke-influenced days. NO_x has been previously
710 seen to have inconsistent enhancement during smoke events (Buysse et al., 2019).

711 The value of θ is seen to increase in the last weeks of the campaign driven by reduced
712 tOHR_{VOC} . Figure S7 compares the VOC contribution to OH reactivity in the days before the air
713 temperature in SLC dropped on September 9th, and those days after. tOHR_{VOC} was lower in the
714 period after the temperature dropped largely due to reduced biogenic terpenes and formaldehyde
715 levels from decreased biological processes in the plants in and around SLC. In addition,
716 formaldehyde also showed a decreased contribution to OH reactivity once the temperature
717 dropped and its photochemical sources reduced due to the seasonal change, which is also
718 captured in our PMF analysis. OH reactivity from NO_x did not have a change between the two
719 time periods, being 2.2 s^{-1} before and after September 9th.

720 Figure 9 also shows the diurnal patterns of tOHR_{VOC} and $\text{tOHR}_{\text{NO}_x}$, separated by weekday
721 versus weekend. tOHR_{VOC} does not show an apparent weekday vs. weekend effect (Section 3.1).
722 NO_x , on the other hand, shows greatly decreased OH reactivity on the weekend compared to the
723 weekdays, likely because decreased traffic and commuters on the weekends led to lower NO_x
724 emissions that have been widely reported in large urban areas including SLC (i.e., Kuprov et al.,
725 2014; Valin et al., 2014). The diurnal patterns for the ratio θ of $\text{tOHR}_{\text{NO}_x}/\text{tOHR}_{\text{VOC}}$ on weekends
726 and weekdays are also plotted. The ratio θ was higher by a factor of 2 in the early daytime hours
727 on weekdays, and larger by a factor of around 1 to 1.5 during afternoon hours, directly
728 corresponding to the difference in $\text{tOHR}_{\text{NO}_x}$ on weekends and weekdays. Future reduction of
729 NO_x emissions will lower the θ thus moving the ozone chemistry to the transitional regime
730 before NO_x -limited. However, based on SAMOZA data, that would require a significant NO_x
731 emission reduction, supporting our modeling study based on the same measurements (Ninneman
732 et al., 2023). Nevertheless, weekends in SLC would likely first show such photochemical regime
733 changes under possible reduced future emissions.

734 One caveat of this analysis is the number of VOCs included in the calculation. We
735 include 35 measured VOCs in SLC, while previous analyses such as Permar et al. (2023)
736 reported that at least 154 VOCs are present in urban air (Boise, ID). Assuming the two cities
737 have similar VOC composition, the results of Permar et al. (2023) would suggest that the 35
738 VOCs included here account for an estimated 60% of tOHR_{VOC} for smoke-free urban air in that
739 study, and roughly 55% of the tOHR_{VOC} in smoke-impacted urban air. Therefore, adding other
740 unreported VOCs in SLC would likely raise the value of tOHR_{VOC} by a factor of 2-3 and thus
741 lower θ by the same magnitude. This would not, however, be a great enough change to change
742 the regime from VOC-limited to transitional or NO_x -limited. Such lower bound estimated θ
743 remains above (though becoming closer to) 0.2 for each day of the campaign, shown in the
744 dashed line in Figure 9. Further, we verify the robustness of the θ indicator for the ozone
745 production regime by comparison to the ratio of formaldehyde to NO_2 concentration (FNR),
746 proposed as an additional metric for use in urban areas (Sillman, 1995; Martin et al., 2004). The
747 daily average of FNR is shown in Figure S8. This metric has previously been used in urban areas
748 with the thresholds showing that ozone production is NO_x -limited when $\text{FNR} > 2$, transitional
749 when $1 < \text{FNR} < 2$, and VOC-limited when $\text{FNR} < 1$ (Duncan et al., 2010; Jin & Holloway, 2015; Li

750 et al., 2021). Most days indeed had $FNR < 1$, with decreasing values in the final weeks of the
751 campaign and becoming more VOC-limited, broadly consistent with the θ metric. Nine days had
752 FNR values between 1 and 2 and therefore showed a transitional regime, consistent with the
753 lower bound estimated θ , though the exact thresholds of θ or FNR could be dependent on local
754 chemical regimes and vary by cities (i.e., Jin et al., 2020; Sourì et al., 2020) .

755 **4 Conclusions**

756 We present measurements of volatile organic compounds (VOCs) and other tracers taken
757 in Salt Lake City, Utah in August-September of 2022 during the SAMOZA field campaign. 35
758 VOCs were measured with online PTR-ToF-MS and offline by DNPH cartridge samples that
759 were subsequently analyzed by HPLC. The total measured VOCs in SLC averaged 32 ± 24 ppb,
760 and the total calculated VOC OH reactivity averaged 3.7 ± 3.0 s⁻¹. Those VOCs with strong
761 photochemical and biogenic sources such as formaldehyde and isoprene showed a 35-50%
762 decrease in concentration after September 9th, 2022, when daily average air temperatures
763 dropped by $\sim 10^\circ\text{C}$ in SLC. Diurnal patterns of most VOCs were typical of urban atmosphere,
764 with high peak mixing ratios in the morning and increasing at night. Methanol and ethanol were
765 the most abundant VOCs present. Terpenes were the largest contributors to OH reactivity in
766 SLC. Contradicted to NO_x, we do not observe an apparent weekday vs. weekend effect in VOC
767 abundance.

768 A PMF solution with 5 factors was used to further analyze VOC sources. Secondary and
769 biogenic, traffic, and solvent uses (or VCPs) from personal care products and industrial sources
770 were found to be important VOC sources in SLC during SAMOZA. Traffic and VCPs had
771 roughly equivalent contributions to anthropogenic VOCs, in agreement with the 2020 NEI.
772 Monoterpenes were found to have a more significant anthropogenic source than biogenic. Source
773 tracer analysis showed high correlations between CO and aromatics, suggesting an important
774 role in traffic emission, similar to the PMF results. Photochemical products like formaldehyde
775 and biogenic VOCs like isoprene had two populations of correlation with CO, indicating some
776 anthropogenic influence. The PMF analysis was used to find a semi-quantitative estimate of the
777 contributions to anthropogenic VOCs in SLC, however further measurements with spatial
778 information are needed to provide more definitive answers.

779 10 out of 61 days during the SAMOZA field campaign were characterized as smoke-
780 influenced. Smoke-free days showed lower abundances of most measured VOCs and thus lower
781 OH reactivity from VOCs. Hazardous air pollutants (HAPs) and other VOCs increased by 45-
782 217% due to the presence of aged smoke from regional wildfires during SAMOZA. The total OH
783 reactivity of VOCs increased by 53% on smoke-influenced days, but the ozone production
784 regime did not transition because of this increase likely due to an enhanced NO_x level during the
785 same period.

786 The ozone photochemical regime was indicated using the ratio between total OH
787 reactivity from NO_x and that from VOCs (or $\theta = tOHR_{NO_x}/tOHR_{VOC}$). The chemical regime was
788 found to be VOC-limited or close to the transitional regime throughout the campaign. The θ
789 increased as temperatures dropped towards the end of the campaign due to the seasonal VOC
790 reductions from biogenic emissions and photochemical sources. $tOHR_{NO_x}$ was much higher on
791 weekdays compared to weekends while $tOHR_{VOC}$ was similar between the two periods. These
792 result in θ being a factor of 1-2 higher on weekdays compared to weekends, thus weekends in
793 SLC may show the first observations of any photochemical regime changes under possible
794 reduced future emissions.

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806 **Open Research**

807 Final SAMOZA data has been archived and can be found at the University of Washington

808 ResearchWorks archive: <http://hdl.handle.net/1773/50049>.

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